

Third assessment of climate change for the Baltic Sea region (BACC III)

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Kommentiert [m1]: In blue auxiliary text from EN CLIME fact sheets.

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Kommentiert [m5]: Not confirmed yet

Kommentiert [m6]: Not confirmed yet

Kommentiert [m7]: Not confirmed yet

Kommentiert [m8]: Not confirmed yet

Abstract. Please use only the styles of this template (MS title, Authors, Affiliations, Correspondence, Normal for your text, and Headings 1–3). Figure 1 uses the style Caption and Fig. 1 is placed at the end of the manuscript. The same is applied to tables (Aman et al., 2014; Aman and Bman, 2015)

5 1 Introduction

1.1 Overview

General introduction, changing climate and knowledge needs for society, policymakers and stakeholders, IPCC and needs for regional information

Focus: Atmosphere, hydrosphere, cryosphere, lithosphere and biosphere

10 Not anthroposphere

1.2 Baltic Sea Region characteristics

Atmosphere:

The climate of the Baltic Sea region is strongly influenced by the large-scale atmospheric circulation (e.g. Andersson, 2002; Tinz, 1996; Meier and Kauker, 2003; Omstedt and Chen, 2001; Zorita and Laine, 2000; Lehmann et al., 2002), in particular the North Atlantic Oscillation (NAO) and blocking and on longer time scales the Atlantic Multidecadal Oscillation (AMO).

15 The NAO is the dominant mode of near-surface pressure variability over the North Atlantic and its impact is strongest during winter (Hurrell et al., 2003), when it accounts for almost one-third of the sea level pressure (SLP) variance. During the positive (negative) phase of the NAO the Icelandic Low and Azores High pressure systems are enhanced (reduced), leading to a stronger (weaker) than normal westerly flow (Hurrell, 1995). For the Baltic Sea region the positive phase of the NAO is related to mild and wet winters and increased storminess (Hurrell et al., 2003).

Atmospheric blocking occurs when persistent high pressure systems interrupt the normally westerly flow over the middle and high latitudes, like e.g. the North Atlantic. This is also frequently observed in the Baltic Sea region. Due to the persistence of blocking events they are often responsible for extreme weather events (Rex, 1950a; Rex, 1950b).

25 The AMO describes fluctuations in North Atlantic sea surface temperature (SST) with a period of 60-90 years (Knight et al., 2006). Thus in the 150-year instrumental record only a few distinct phases have been observed. However, a recent model study suggests that variations in the AMO may have an impact on the precipitation over the Baltic Sea region (Börgel et al., 2018).

Kommentiert [m9]: Introduction into the region and differences between the Baltic Sea and other coastal seas

Precipitation is water falling to the ground. It can take various forms including melted (e.g. rainfall and drizzle) or frozen (e.g. snowfall and hail) or in mixed forms involving both snow and rain (e.g. sleet). Precipitation is measured in mm of melted water over a certain time interval that could include one or several precipitation events.

- 5 Precipitation is strongly linked to other parameters describing the water cycle. The amount of water in the air is one of the most important factors implying that the history of an air mass including previous evaporation and precipitation events is important. As the amount of water that can be held in air depends on temperatures, precipitation has some temperature dependency (generally more precipitation in summer than in winter in large parts of the Baltic Sea region). Precipitation is also strongly modified by orographic features implying that the large-scale circulation of the atmosphere including wind direction and vertical stability are important factors.

Precipitation is a key feature in determining soil moisture conditions, runoff and discharge. The impact on soil moisture can represent a strong feedback mechanism as dry conditions lead to less evaporation and thereby less precipitation etc.

- 15 Precipitation is one of the key parameters in the water cycle, it can be stratiform, convective and orographic and is depending on atmospheric circulation patterns, convection and accessibility to water.

Land:

- 20 The magnitude of water flow in a river is the result of various complex hydrological processes including *precipitation*, *evapotranspiration*, *infiltration* and *storage* (e.g. in the form of snow, soil moisture, and sub-surface and groundwater storage). Therefore, explaining changes in streamflow requires an understanding of these parameters of which precipitation is often the pivotal one in the cool climate of the Baltic Sea region.

- 25 River runoff and net precipitation (precipitation minus evaporation) over the sea surface are dominant drivers of Baltic Sea salinity explaining together with the limited water exchange with the North Sea the large gradient in sea surface salinity between about 20 g kg⁻¹ in Kattegat and 2 g kg⁻¹ in the Bothnian Bay (REF). Net precipitation amounts to about 10% of the total river runoff (e.g., REF, Meier and Döscher, 2002).

- 30 Ocean:

Salinity:

Due to freshwater supply from the Baltic Sea catchment area and due to the limited water exchange with the world ocean, surface salinity varies from $> 20 \text{ g kg}^{-1}$ in Kattegat to $< 2 \text{ g kg}^{-1}$ in the Bothnian Bay. The dynamics of the Baltic Sea is characterized by a two layer system because of a pronounced, perennial vertical gradient in salinity.

5 Meteorologically driven large saltwater inflows (so-called Major Baltic Inflows, MBIs) sporadically renew the deeper parts of the Baltic Sea with saline, oxygen rich water. Hence, MBIs are the only process that effectively ventilate the deep water. However, MBIs are estimated to contribute to the total salt important by only 20%.

Stratification:

10 Stratification denotes the vertical layering of water bodies according to different water densities mainly determined by layer specific water temperature and salt concentrations. Stratification controls vertical transports in the water column, e.g. the downward flow of dissolved oxygen from the sea surface into the deep layers. A number of parameters exist to characterize stratification. Pycnocline commonly denotes the strongest density gradient in the vertical water column. Corresponding definitions exist for the temperature gradient (thermocline) and salinity gradient (halocline) which highlight the governing parameters for the density layering.

The forces of the winds (i.e. wind stress working at the air water interface) can potentially homogenize the water column fully (some shallow water regions) or partly (deep water regions) and thus influence stratification.

20 In the Baltic Sea a pronounced halocline persists over the year between 60-80 meters in most regions. During the warm season a thermocline develops at much shallower depths (10-20 meter).

Theoretical considerations imply that increased freshwater supply (precipitation minus evaporation, snow melt) over the Baltic Sea drainage basin accompanied by the supply of deep salt rich waters from the North Sea as well as warming of the surface layer would favor stronger stratification.

Sea level:

30 Sea level is traditionally measured with a scale on the harbor wall. The earliest scales from the 18th century were marks on large, prominent boulders along the shoreline of the Baltic Sea (Ekman, 2009). Sea level changes when water is added to the global ocean, when it expands by warming or when the land is rising to which the scale is attached. These are the three most important contributions to long-term changes in sea level in the Baltic Sea (BACC I, Hünicke et al., 2015). Changes in oceanic and atmospheric circulation cause variations of sea levels in the Baltic Sea on seasonal to decadal time scales (Chen and Omstedt, 2005). Winter sea level in the Baltic Sea is usually higher than summer sea level (Samuelsson and Stigebrandt, 1996). And mild winters with stronger than average winds show higher sea levels than severe winters (Andersson, 2002, Karabil et

al., 2018). During the cold seasons from fall to spring the risk is higher for storms to produce storm surges along the coasts of the Baltic Sea (Lehmann et al., 2011). Storm surges are water masses pushed against the coasts by the wind of atmospheric low pressure systems. The lower atmospheric pressure in a storm can add to the amplitude of the storm surge. Storms also excite oscillations in the Baltic Sea that are highest at the ends of the different basins and lowest in the Baltic Proper (Weisse and Weidemann, 2017). Waves that are generated during storms can add to extreme sea levels measured along the coast (Eelsalu et al., 2014). The amplitude of tides is relatively small in the Baltic Sea and adds to extreme sea levels only in the Kattegat and Skagerrak.

The mean sea level of the global ocean is closely related to the volume of the ocean. The volume is changing due to geologic activity and to the addition of water that was previously frozen on land like the Antarctic and Greenland ice sheets and mountain glaciers. Melting rates of glaciers are sensitive to changes in air temperature, snow cover and the resulting brightness. Ice sheets in polar regions are affected additionally by contact with ocean water melting the ice from below, which can cause instabilities in the integrity and the flow of the ice sheets. Ocean temperatures affect the mean sea level because warmer water is lighter and takes up more volume than colder water. The continuous redistribution of heat from the tropics to the polar regions establishes the ocean circulation systems that form a sea surface relief along which the ocean currents are flowing. This sea surface relief modulates the global mean sea level by up to half a meter (e.g. Yin et al., 2009). The mean sea level in the Baltic Sea depends on the strength of the ocean circulation in the Atlantic (Yin et al., 2009, 2010, Yin, 2012, Balmaseda et al., 2013b). Temperature and freshwater distribution in the ocean leave their imprint on height of the sea surface and cause it to vary on decadal time scales (Balmaseda et al., 2013a). Atmospheric winds are another important driver of sea level variability on time scales from decades down to a couple of hours during a storm surge.

Sedimentation:

Sediment transport is triggered mainly by currents in marine environment, by waves in the nearshore and by wind in subaerial coastal environment. Its direct consequence is a gradual change of the earth surface landform, leading to erosion or accretion.

Short-term and small-scale sediment transport is strongly hinged on a variety of local state variables including wind velocity and direction, water level, waves as well as the antecedent state of the system, while long-term and large-scale sediment transport and coastal erosion are primarily controlled by sediment supply modulated by large-scale processes, notably mean sea level, storms, the regional wind and wave pattern, and engineering structures.

Due to a combined effect of isostatic adjustment and eustatic sea level change, the coastline change of the Baltic Sea is characterized by a North-South gradient from an uplift of max. 9 mm/yr in the North to a subsidence of min. -2 mm/yr in the south since the onset of the Holocene. The subsiding southern Baltic Sea coast is characterized by a series of barrier islands

and sandy dunes connected with soft moraine cliffs. The composition of soft, mobile sediments makes the southern Baltic Sea coast extremely vulnerable to wind-wave induced transport and erosion.

Most coastline erosion along the southern Baltic Sea is caused either by storms or human-induced depletion of sediment supply (e.g. side effect of engineering structures).

5

Oxygen concentration and hypoxia:

Dissolved oxygen concentration in the water column is controlled by physical transport supply (advection and diffusion) and biological oxygen demand for oxidation of organic matter. Due to limited ventilation largely incapable to meet the oxygen demand by elevated concentrations of organic matter in the water column and sediments (eutrophication), the Baltic Sea deeps suffer from deoxygenation and hypoxia. Hypoxic area is defined as the extent of bottom water with oxygen concentrations below a threshold such as 2 mL O₂ L⁻¹. Hypoxia is characterized by the lack of higher forms of life.

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1.3 Global climate change

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1.4 The BACC process

What is new in BACC III? BACC III is an update of BACC II, assessment of knowledge from literature published after BACC II (focus period 2013-2019), ESD peer-reviewed journal, independent editors, independent scientific reviewers, part of BEAR (9 review articles), common figures of atmospheric, hydrological and oceanographic projections based on published climate simulations, common uncertainty analysis

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1.5 Summary of BACC I and II key messages

1.5.1 Past climate change

Long-term changes on paleoclimate timescales

1.5.2 Recent climate change

Climate change observed during the instrumental period

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1.5.3 Future climate change

Projections based upon IPCC (2007)

2 Methods

2.1 Assessment of literature

Assessment of scientific peer-reviewed literature and reports, statistics of references

2.2 Climate model data

5 Published model data based upon IPCC (2013),
uncoupled simulations with several RCMs and GCMs (Ole, Erik) and coupled simulations with one RCM and 8 GCMs
(Matthias, Christian, Markus), RCP 2.6, 4.5 and 8.5

10 An additional ensemble is available that has been produced with a coupled RCM. The RCM RCA4-NEMO has been introduced
by Wang et al., 2015. Gröger et al., 2015, 2019 and Dieterich et al. 2019 have validated different aspects of the ensemble
discussed here. The atmosphere component RCA4 is run at a resolution of 0.22 [degrees] and 40 levels in the EURO-CORDEX
domain. Coupled to it is the North Sea-Baltic Sea model NEMO-Nordic at a resolution of 2 nautical miles and 56 levels. The
two components of the RCM are coupled by sending sea level pressure, energy, mass and momentum fluxes every 3 hours
15 from the atmosphere to the ocean. Vice versa, the atmosphere receives at the same frequency sea and ice surface temperatures
and the sea ice fraction and albedo.

Table 2: Ensemble members of three RCP scenarios with the coupled RCM.

| | RCP8.5 | RCP4.5 | RCP2.6 |
|--------------|-------------|-------------|-------------|
| EC-EARTH | 1961 - 2099 | 1961 - 2099 | 1961 - 2099 |
| MPI-ESM-LR | 1961 - 2099 | 1961 - 2099 | 1961 - 2099 |
| HadGEM2-ES | 1961 - 2099 | 1961 - 2099 | 1961 - 2099 |
| GFDL-ESM2M | 1961 - 2099 | 1961 - 2099 | 1961 - 2099 |
| IPSL-CM5A-MR | 1961 - 2099 | 1961 - 2099 | 1961 - 2099 |
| CanESM2 | 1961 - 2099 | 1961 - 2099 | 1961 - 2099 |
| NorESM1-M | 1961 - 2099 | 1961 - 2099 | 1961 - 2099 |
| MIROC5 | 1961 - 2099 | 1961 - 2099 | 1961 - 2099 |

20 This RCM has been used to downscale eight different GCMs with three RCPs each (Table 2). Surface variables of the
atmosphere component are saved at hourly to 6-hourly frequency to allow for an analysis of means and extremes in present
and future climates.

2.3 Uncertainty estimates

IPCC terminology, methods applied in this study

3 Current state of knowledge

5 3.1 Past climate change

Knowledge gained from proxy data and paleoclimate modeling

3.2 Present climate change

10 The past 200 years (instrumental records, model based reconstructions and reanalyses), changes in mean and extremes, trends and decadal/multi-decadal variability

3.2.1 Atmosphere

Large-scale circulation (teleconnection patterns), surface pressure, wind, surface atmosphere temperature, precipitation, cloudiness and solar radiation

3.2.1.1 Large-scale circulation

15 NAO:

While the NAO exhibited a positive trend from the 1960s to the 1990s it has returned to lower values in the early 2000s with exceptionally low anomalies in the winters of 2009/2010 and 2010/2011, which considerably weakened the positive trend.

Blocking:

20 While some studies find an eastward shift of blocking events over the North Atlantic (Davini et al., 2012; Croci-Maspoli et al., 2007) and increase in blocking duration over the Northern Hemisphere since about 1990 (Mokhov et al., 2013), there is low confidence in these changes due to methodological differences between studies (IPCC, 2013).

AMO:

25 The AMO has been warming from the late 1970s to the 2000s as part of its natural variability and has since remained in a warm state. Natural fluctuations in the AMO over the coming few decades will likely influence regional climates, like e.g. the Baltic Sea region, at least as strongly as human-induced changes (IPCC, 2013).

increasing linear trend of the number of deep cyclones for the period 1950-2010 (Lehmann et al., 2016)

Kommentiert [m10]: But there is no long-term trend since 1850.

3.2.1.2 Air temperature

Air temperature introduction: Air temperature shows the clearest response to the increased green-house effect. Changes in temperature extremes may influence human activity much more than changes in average temperature.

5 **Air temperature:** A significant surface air temperature increase in the Baltic Sea region during the last century was detected (BACC Author Team 2008; 2014; Rutgersson et al., 2014). The temperature increase is not monotonous but accompanied by large (multi-) decadal variations dividing the 20th century into 3 main phases: (1) warming in the beginning of the century until the 1930s; (2) cooling until 1960s; and (3) another distinct warming during the last decades of the time series. Linear trends of the annual mean temperature anomalies during 1871–2013 were 0.10 K decade⁻¹ north of 60°N and 0.08 K decade⁻¹ south of 60°N in the Baltic Sea region. This is larger than the global mean temperature trend. There are large variations, in particularly during winter, but the warming is seen for all seasons (being largest during spring). These changes are also resulting in seasonality changes: the length of the growing season has increased, whereas the length of the cold season has decreased. The number of days by which autumn and winter are delayed differs from south to north and east to west, but as an example in Tartu, Estonia, the number of deep winter days (with snow cover) has decreased by 29 d over the past century while the growing season has increased by 13 d in this period (Kull et al. 2008).

Air temperature extremes: The duration of extremely mild periods has increased significantly in winter, while the number of heat waves has increased in summer as well as during the year as a whole. A general increase has been observed for the annual numbers of days with daily maximum temperature above both 25 and 30 °C, and a decrease in the length of the frost season and in the annual number of frost days.

3.2.1.3 Solar radiation

Solar radiation introduction:

Total cloudiness consists of clouds on all levels (low, medium and high) and is related to the general circulation as well as the water cycle. Solar radiation is to a large extent depending on the cloudiness (amount and type of clouds), but also to atmospheric aerosols. Atmospheric aerosols affect solar radiation under clear skies directly and through interaction with clouds indirectly.

Solar radiation: Multidecadal variations, known as “dimming” and “brightening” have been observed both in Europe and many other parts of the world, e.g. Wild et al. 2005, Wild et al. 2012, Wild et al. 2017 (especially on the northern hemisphere). No long-term measurements over oceans. However, (aerosol-induced) multidecadal variations in surface solar radiation could be expected also over oceans (Wild et al. 2016).

Satellite data records of trends in cloudiness since the 1980s disagree over many areas but there is some consistency about a decline in cloudiness over the Baltic area (Karlsson and Devasthale, 2018 (fig 4)).

3.2.1.4 Precipitation

Anna:

- 5 Precipitation is highly variable, and the amount has varied between regions and seasons, with both increasing and decreasing precipitation and **no general trend for the entire region**, however areas have experienced both increasing and decreasing trends.

Kommentiert [m11]: Disensus?

Erik:

On average, annual mean precipitation has generally increased over most of the Baltic Sea region over the 20th Century.

Kommentiert [m12]: Disensus?

- 10 Differences between different parts of the region and between seasons are prominent.

Precipitation extremes:

Anna:

- 15 An increase in extreme precipitation and duration of the lengths of wet periods. Also the length of dry periods have been seen in some areas.

Decreased summer precipitation and increased length of dry periods is closely linked to increased risks of forest fires and limitations in accessibility to drinking water. It would also cause problems for agriculture. More intense precipitation increases the risks of flooding, if drainage systems are not adapted.

20

Erik:

Precipitation increase in northern Europe is generally associated with an increase in the frequency and intensity of extreme precipitation events. Observed changes include increasing intensity and/or frequency of intense precipitation events, changes in duration of wet and dry spells.

- 25 **3.2.1.5 Wind**

Wind

Wind extremes

3.2.1.6 Atmospheric chemistry

3.2.2 Land

River runoff, nutrient and carbon loads, river ice, lakes, lake ice, snow, permafrost, terrestrial and freshwater ecosystems

3.2.2.1 River discharge

- 5 Run-off to the Baltic Sea appears to be strongly linked to temperature, wind and rotational circulation components in the northern region and Gulf of Finland. In contrast, run-off in the southern region is more associated with the strength and torque of the cyclonic or anticyclonic pressure systems. (Hansson et al. 2011)

Although decadal and regional variability is large, no *statistically significant* long-term change has been detected in the reconstructed total river runoff to the Baltic Sea over the past 500 years.

- 10 As a whole, over the past 500 years, the total river runoff to the Baltic Sea has decreased slightly in response to the rise in temperature; at a rate of 3%, or $450 \text{ m}^3 \text{ s}^{-1}$, per 1°C (Hansson et al. 2011).

The observed temperature increase has clearly affected stream flow in the Nordic countries. These changes correspond well to the projected consequences of a continued rise in global temperature, whereas the impacts of both the observed and projected changes in precipitation on stream flow are unclear (Hisdal et al. 2010).

- 15 In the Baltic States (Lithuania, Latvia and Estonia) in general, changes in stream flow over the 20th century show a redistribution of runoff throughout the year: with a significant increase in winter river discharge and a tendency for decreasing spring floods (Reihan et al. 2007).

river flow to the Bothnian Bay show a statistically significant positive trend during 1921-2004 (Kniebusch et al., 2019)

- 20 Stahl et al. (2010)

Kniebusch et al. (2019b)

No statistically significant linear trends in annual river discharge to the sea has been detected. Winter flows have increased due to temperature while spring flows have decreased (BACC II)

- 25 For the period 1850-2008, the total river runoff from the Baltic Sea catchment area reconstructed from observations (Hansson et al., 2011a; Cyberski and Wroblewski, 2000; Mikulski, 1986; Bergström and Carlsson, 1994) and hydrological model results (Graham, 1999) shows no statistically significant trend but a pronounced multi-decadal variability with a period of about 30 years (Meier et al., 2018b).

3.2.2.2 Lakes

- 30 3.2.2.3 Terrestrial biogeochemistry

Nutrient loads

Kommentiert [m13]: This number is rather uncertain. The reconstructed runoff by Hansson et al. 2010 underestimates interannual variability considerably. Multidecadal variability is almost missing

Kommentiert [m14]: Not only in the Baltic States. Isn't the winter mean trend of the total river flow statistically significant as well? Check! Check!

In the reference scenario, nutrient loads represent the average loads of the period 2010-2012. The high or worst case scenario assumes changes caused by a 'fossil-fuelled development' scenario coupled to increasing river runoff. Changes in nitrogen and phosphorus loads were calculated from the regional assumptions, e.g., on population growth, changes in agricultural practices such as land and fertilizer use and expansion of sewage water treatment plants.

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PLC reports statistically significant reductions in riverine nutrient loads to the sea relative to the 1997 to 2003 reference period (PLC6).

These reductions are not attributed to climate.

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DOC inputs to the sea have increased over the past century but the cause is not known (BACC II).

3.2.2.4 Terrestrial ecosystems and biodiversity

Land cover

Agriculture and forest

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Birds

3.2.3 Cryosphere

3.2.3.1 Snow

3.2.3.2 Glaciers

3.2.3.3 Permafrost

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3.2.3.4 Sea ice

Looking back a little bit more than 100 years, the ice winters have become milder, the maximum ice extent has decreased and the ice season is shorter. Also indexes based on the total ice volume of the winter show a decreasing trend. Severe ice winters can still happen nowadays, but the possibility therefore has decreased.

25

The maximum ice extent in the Baltic Sea is changing between 50 and 400 km² from year to year. During mild winters just the Bothnian Bay is ice covered while in severe winter the whole Baltic Sea can be ice covered like in the ice winter 1986.

3.2.3.5 Cryospheric ecosystems and biodiversity

Kommentiert [m15]: More information is available

3.2.4 Ocean and sediment

Water temperature, salinity, stratification, currents, water exchange, sea ice, sea level, wind waves, sediment transports, marine ecosystems (including coastal ecosystems, biogeochemistry), acidification,

3.2.4.1 Water temperature

- 5 The temperature rises fastest at the surface. With time the heat spreads downward through different processes and eventually the whole water column warms up. This has consequences for the stratification (EN-CLIME stratification), eutrophication (EN-CLIME nutrient cycle) and sea level rise (EN-CLIME sea level).

10 The marginal seas around the globe have warmed more than the average over the global ocean (Belkin, 2009). The Baltic Sea has warmed during the second half of the past century, although the interannual variability is high (Kniebusch et al., 2019a). During the period 1982 to 2006 the SST increased by 1.35 C in the Baltic Sea. No other sea has warmed up that much. Fonselius and Valderrama (2003). The accelerated warming is partly explained by the phase of the AMO (Kniebusch et al., 2019a).

3.2.4.2 Salinity and saltwater inflows

- 15 New research results contributed to the understanding of the role of natural variability and climate change for salinity changes in the Baltic Sea (e.g., Börgel et al., 2018; Kniebusch et al., 2019b; Lehmann and Post, 2015; Lehmann et al., 2016; Mohrholz, 2018; Schimanke and Meier, 2016). A new reconstruction of saltwater inflows for the period 1887–2017 showed that the apparent decline in frequency and intensity of MBIs in the reconstructed record by Fischer and Matthäus (1996) since 1976 was based on inhomogeneous data (Mohrholz, 2018). Hence, we can conclude that there are no statistically significant trends
- 20 in salinity (Fonselius and Valderrama, 2013; Meier et al., 2018b; Kniebusch et al., 2019b), stratification (Meier and Kauker, 2003), river flow (Meier and Kauker, 2003a; REF), large volume changes (LVC) (Lehmann and Post, 2015) and MBIs (Mohrholz, 2018) on centennial time scales since about 1900. However, salinity, river flow and MBIs showed a pronounced multi-decadal variability with a period of about 30 years (Winsor et al., 2001; 2003; Gailiūšis et al., 2011; Mohrholz, 2018). Part of this variability is, for instance, the stagnation period during 1983-1992 without MBI and with decreasing salinity
- 25 (Nehring and Matthäus, 1991).

According to model results, multi-decadal variations in runoff explained about 50% of the long-term variability of volume averaged salinity of the Baltic Sea (Meier and Kauker, 2003a). Radtke et al. (2019) found that the direct dilution effect was responsible for about one fourth of the multidecadal variability only and that the impact of vertical turbulent mixing is small.

- 30 Salt water inflows contribute to the multidecadal variability in salinity as well, in particular of the bottom layer salinity. Further, model results suggest that decreasing salinity over ten years appear approximately once per century on average and belongs to

the natural variability of the system (Schimanke and Meier, 2016). On longer time scales, Baltic Sea salinity is under the influence of the Atlantic Multidecadal Oscillation with a period of about 60-90 years (Börgel et al., 2018). Probably part of this long-term natural variability, since about the 1980s observed surface and bottom salinities decreased and increased, respectively (Vuorinen et al., 2015; Liblik and Lips, 2019).

5

Since about the 1970s, the mean seasonal cycle of the total river flow has changed with increasing and decreasing runoff during winter and summer, respectively (REF; Meier and Kauker, 2003a). These changes might be explained by river regulation of large rivers in the North (Matthäus and Schinke, 1999) and systematic changes in precipitation patterns due to warming in the Baltic Sea region (REF). However, as the change in seasonality does not affect the total discharge trend and as there is no statistically significant trend in saltwater inflows on centennial time scale (Mohrholz, 2018), changes in salinity are regional only. In Baltic Sea reconstructions and local observations, Kniebusch et al. (2019b) found a statistically significant positive trend in the North-South gradient of sea surface salinity for 1900-2008. This change is mainly attributed to increased river runoff from the northernmost catchment indicating a footprint of the anthropogenic impact on salinity.

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3.2.4.3 Stratification

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Direct consequences of increasing stratification is that mixing between well ventilated surface waters and deep waters weakens. This makes the Baltic Sea more vulnerable against deoxygenation of bottom waters (haline stratification). An increase in seasonal thermal stratification can additionally lower the vertical nutrient transport from deeper layers to the euphotic zone thereby limiting nutrient supply and potentially facilitate cyanobacteria blooms in the Baltic Sea.

3.2.4.4 Overturning circulation

20 3.2.4.5 Sea level

Global mean sea level rise is measured at 1.5 mm/year during the 20st century. It is estimated at a rate of 2.8 mm/year during the period since satellites are measuring (1993 to 2010). The rate of global mean sea level rise is increasing [high confidence]. During the second half of the past century thermal expansion of sea water and the addition of melt water from the Antarctic and Greenland ice sheets and from glaciers have contributed about equally to the process [high confidence] (Church et al., 2013). Sea level rise (relative to the ellipsoid) in the Baltic Sea in the past 50 years is estimated between 1 and 3 mm/year (Milne et al., 2001, Hill et al., 2010, Richter et al., 2011, BACC I, BACC II). For the period 1886-2017, Swedish Baltic Sea records showed an increase of about 24 cm corresponding to 1.8 mm/year (Source: SMHI).

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Sea level extremes:

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Return levels with corresponding return periods is a concept to assess the amplitude and the probability of extreme events. A storm surge with a 100-year return period will occur on average once in 100 years. The amplitude of the event depends on the

Kommentiert [m16]: Uncertainty range from the literature. What are the latest numbers? Is there an acceleration since 1980?

location in the Baltic Sea. Storm surges at the end of the basins furthest away from the Baltic Proper are higher than those in the center of the Baltic Sea [high confidence] (Meier et al., 2004, Wolski et al., 2014, Dieterich et al., 2019). No long-term trend has been found so far that points to an increase of extreme sea levels in the Baltic Sea (EN-CLIME wind, BACC II). There are however studies (e.g. Suursaar and Sooäär, 2007) that confirm extreme sea level e.g. in the Gulf of Finland to be very sensitive to the position of the storm tracks.

Kommentiert [m17]: Subtracting the mean sea level rise?

Kommentiert [m18]: What does this mean? Will the projected northward shift of the storm track have an impact on extremes? You should also add the results by Lang and Mikolajewicz (2019).

No changes during different climate conditions (Lang and Mikolajewicz, 2019)

3.2.4.6 Waves

10 3.2.4.7 Sedimentation

Mean erosional rate along Latvia and Lithuania for the sandy coasts and soft cliffs are 1–2 m/yr and 0.5-0.6 m/yr in the latter half of the twentieth century, respectively. Since 1980s erosional rates of certain sections have been enhanced to 1.5–4 m/yr;

15 Mean erosional rate along Lithuania and Russia (Kaliningrad) for the sandy coast and cliff are 0.5-0.8 m/yr and 1-1.5 m/yr, respectively, and rates have increased over the past decade.

The Polish coast is mainly formed of soft sandy sediments, with an average rate of retreat of 0.5–1.5 m year⁻¹. Coastal towns which experienced erosion of 0.3–0.7 m/yr now have a nourished beach.

20 Mean erosion rate along the German sandy coast is 0.4 m/yr.

In southern Sweden, the soft moraine cliffs have retreated 1–1.5 m year⁻¹ over the past 150 years.

Erosion rates along the dune coasts in the Kattegat are within 2 m year⁻¹.

25

Coastal erosion has the following direct consequences:

1. loss of coastal lands;
2. loss of coastal resilience;
3. loss of valuable natural habitats;
- 30 4. loss of economic value and private property;
5. increasing cost to society in terms of coastal protection.

Coastal erosion extremes:

Many sandy beaches along the Gulf of Finland have recently been severely damaged by frequent storm surges, despite extensive protective measures.

5

Soft cliffs in Latvia are eroded 3–6 m/yr after each storm, with a maximum of up to 20–30 m/yr at local sites.

Extreme erosional rate along Lithuania and Russia (Kaliningrad) for cliff are 10 m/yr.

10 Poland: recent erosion after each storm surge reaches 3–6 m/yr;

Germany: erosion by storms has reached 3 m/yr at some local coastal sections

Maximum erosion rates along the dune coasts in the Kattegat are ~2 m/yr.

15

3.2.4.8 Marine carbonate and biogeochemistry

Air-sea exchange of CO₂:

20 There is clear seasonal pattern in the partial pressure of CO₂ (pCO₂) in the surface Baltic Sea. It is controlled by the biologically driven processes (organic matter production and remineralization) as well as changes in the mixed layer depth.

In winter the surface Baltic Sea is oversaturated with CO₂, while in the productive periods pCO₂ goes down below the atmospheric level with 2 clear minima: during spring bloom and N₂-fixing cyanobacteria bloom.

25

The difference between seawater pCO₂ and atmospheric pCO₂ during the productive periods has increased due to the eutrophication and development of cyanobacteria blooms.

During spring bloom it changed from about 50 μatm (Buch, 1945) to ca. 250 μatm (BACC II), while during summer from about 40 μatm (Buch, 1945) to 300 μatm (BACC II).

30

The pCO₂ seasonal cycle in the surface water controls the annual CO₂ exchange through the air/sea interface.

A_T is increasing in the Baltic Sea. The highest A_T trend was found in the Gulf of Bothnia ($7.0 \mu\text{mol kg}^{-1} \text{yr}^{-1}$), followed by $3.4 \mu\text{mol kg}^{-1} \text{yr}^{-1}$ in the central Baltic whereas no trend could be detected in the Kattegat (Müller et al., 2016).

5 The mean CO_2 concentration is increasing in the surface Baltic Sea due to the increase of CO_2 in the atmosphere. (Schneider et al., 2015, BACC II).

Acidification:

10 According to the thermodynamics of the CO_2 system the increase of 2.0 ppm yr^{-1} in the atmospheric pCO_2 should result in the pH decrease of about 0.02 per decade if the total alkalinity would be constant.

However, the pH trend expected based on the pCO_2 increase could not be definitively identified in the Baltic Sea.

Ocean acidification (pH decrease) is to large extent mitigated in the Baltic Sea by the total alkalinity (A_T) increase (Müller et al., 2016).

15

The highest A_T increase was found in the Gulf of Bothnia ($7.0 \mu\text{mol kg}^{-1} \text{yr}^{-1}$), followed by $3.4 \mu\text{mol kg}^{-1} \text{yr}^{-1}$ in the central Baltic whereas no trend could be detected in the Kattegat.

Other drivers:

20 Enhanced organic matter production (eutrophication) and remineralization causes high amplitude of the seasonal pH changes in the Baltic Sea

Oxygen concentration and hypoxia:

25 Despite the decrease of nutrient loads from land after the 1980s, recently observed oxygen consumption rates are higher than ever observed, limiting the impact of natural ventilation by oxygen-enriched saltwater intrusions in the open Baltic Sea. Improving oxygen conditions have been observed in some coastal systems, where inputs of nutrients and organic matter have been abated. However, hypoxia remains a large problem for many coastal systems, displaying unaltered or even worsening conditions.

30 In the Baltic Sea hypoxia has expanded considerably since the first oxygen measurements became available in 1898. In 2016 the annual maximum extent of hypoxia covered an area of about $70,000 \text{ km}^2$, comparable with the size of Ireland, whereas 150 years ago hypoxia was presumably not existent or at least very small. Hypoxia was mainly caused by accumulation of increasing riverborne nutrient loads and atmospheric deposition. The impacts of other drivers like observed warming and eustatic sea level rise were comparatively smaller but still important.

5 Despite the decrease of riverborne nutrient loads after the 1980s, recently observed oxygen consumption rates are higher than ever observed, limiting the impact of natural ventilation by oxygen-enriched saltwater intrusions in the open Baltic Sea (Meier et al., 2018a).

10 In the Baltic Sea, hypoxia has expanded considerably since the first oxygen measurements became available in 1898. In 2016, the annual maximum extent of hypoxia covered an area of about 70,000 km², comparable with the size of Ireland, whereas 150 years ago hypoxia was presumably not existent or at least very small (Carstensen et al., 2011; Meier et al., 2018b). Hypoxia was mainly caused by accumulation of increasing riverborne nutrient loads and atmospheric deposition (Savchuk, 2010; 2018). The impacts of other drivers like observed warming and eustatic sea level rise were comparatively smaller but still important (Carstensen et al., 2011; Meier et al., 2018b). Hence, also halocline variations had an impact on hypoxia (Väli et al., 2013). Although in some coastal regions, as a consequence of implemented abatements, improved oxygen conditions have been observed, persistent and seasonal hypoxia remains a large problem for many coastal systems (Conley et al., 2011).

15 Future changes:

20 Projected warming and global mean sea level rise may reinforce eutrophication and oxygen depletion in the Baltic Sea by reducing air-sea fluxes and vertical transports of oxygen in the water column, intensifying internal nutrient cycling, and increasing river-borne nutrient loads due to increased river runoff (Meier et al., 2011; 2012a; 2012b). However, the response of deep-water oxygen conditions to changing climate will mainly depend on the nutrient load scenario (Saraiva et al., 2019a; 2019b). In the case of high (low) nutrient loads, the impact of the changing climate would be considerable (negligible). Scenario simulations suggest that the complete implementation of the Baltic Sea Action Plan (BSAP) resulting in required load reductions will lead to a significantly improved ecosystem state of the Baltic Sea irrespective of the driving global climate model (Saraiva et al., 2019b) and regional coupled climate-environmental model (Meier et al., 2018c).

25 For the end of the century (2069-2098), hypoxic area is projected to change only slightly in the ensemble mean under reference (-14 ... -5%) and high (-2 ... +5%) nutrient load scenarios compared to the period 1976-2005 (Saraiva et al., 2019b). In the reference scenario, nutrient loads represent the average loads of the period 2010-2012. The high or worst case scenario assumes changes caused by a 'fossil-fuelled development' scenario coupled to increasing river runoff (Saraiva et al., 2019a). Changes in nitrogen and phosphorus loads were calculated from the regional assumptions, e.g., on population growth, changes in agricultural practices such as land and fertilizer use and expansion of sewage water treatment plants (Zandersen et al., 2019). Under the BSAP scenario, hypoxic area will at the end of the century be reduced by 50 to 60% in the ensemble mean compared to 1976-2005 (Saraiva et al., 2019a).

For the first time, recently uncertainties in projections of the Baltic Sea ecosystem have systematically been assessed (Meier et al., 2018c; 2019). One of the identified larger sources of uncertainty are caused by biases of global and regional climate models, in particular, with respect to GMSL rise and regional water cycling (Meier et al., 2019). The mechanism behind the correlation between large-scale meteorological conditions in the different climate periods and O₂ conditions in the Baltic Sea is not well understood and subject to ongoing research.

Kommentiert [m19]: From Lehmann et al.

Secchi depth:

3.2.4.9 Marine ecosystems and biodiversity

Pelagic habitats (incl. phytoplankton and zooplankton community structure, spring blooms, functional traits etc.)

Spring bloom phenology (timing and length): Observed for a time series 2000-2014 from the Baltic Proper and the Gulf of Finland; consequence is potential mismatch with timing of zooplankton (Groetsch et al., 2016), other drivers: warming and prevalent high pressures

Changes in plankton communities: Observed on the northern Baltic Sea between 1979 and 2011 (Suikkanen et al., 2013), a complex set of drivers including warming, eutrophication and trophic cascades

Summer biomass of cyanobacteria: Observed on the northern Baltic Sea between 1979 and 2011 (Suikkanen et al., 2013), other drivers: warming, decreasing salinity and changes in nutrient ratios

Spring bloom dinoflagellate biomass increase: Observed in some basins (Klais et al., 2013; BACC II, 2015), other drivers: ice free winters favor dinoflagellates

Wasmund (2017), Wasmund et al. (2017)

Benthic habitats (incl. benthic organisms and community structure, functional traits)

Fish

Water birds

Marine mammals

Non indigenous species

Ecotoxicology

Ecosystem function

3.3 Future climate change

3.3.1 Atmosphere

Large-scale circulation (teleconnection patterns), surface pressure, wind, surface atmosphere temperature, precipitation, cloudiness and solar radiation,

- 5 Figures of ensemble mean changes in mean and extreme variables in uncoupled (Ole, Erik) and coupled simulations (Matthias, Christian)

3.3.1.1 Large-scale circulation

NAO:

- 10 In the future, the NAO is very likely to continue to exhibit large natural variations similar to those observed in the past. It is likely to become on average slightly more positive due to an increase in greenhouse gases (GHG) (IPCC, 2013).

Blocking:

There is medium confidence that the frequency of blocking will not increase. However, trends in the intensity and persistence of blocking remain uncertain and therefore also the implications of blocking related changes in the Baltic Sea region (IPCC, 2013).

- 15 AMO:

Based on paleoclimate reconstructions and long model simulations it is unlikely that the AMO will change its behaviour in the future under a changing mean climate (IPCC, 2013).

- 20 consistently across the ensemble a northward shift in the mean summer position of the westerlies at the end of the twenty-first century compared to the twentieth century (Gröger et al., 2019). century. Associated with this is an anomalous precipitation pattern marked by increased rainfall over northern Europe and dryer conditions over the continental central part. In response to these large-scale atmospheric changes, a strong freshening mainly resulting from a higher net precipitation over the year combined with higher annual mean runoff is registered for the Baltic Sea and adjacent seas.

3.3.1.2 Air temperature

- 25 Air temperatures in the Baltic Sea area are projected to increase with time, with the increase generally greater than the corresponding increase in global mean temperature. This is usually the case for land areas, which warm more quickly than sea areas but is also the case for the Baltic Sea region, largely due to the strong winter increase. This winter increase is the result of a positive feedback mechanism involving declining snow and sea-ice cover, leading to even higher temperatures—reduced snow and ice cover will enhance the absorption of sunlight, and so enable greater amounts of heat to be stored in the soil and
30 water (BACCII, 2015).

The increase in winter temperatures are projected to reach 8 degrees C in northern Scandinavia depending on the representative concentration pathway (RCP). The RCP2.6 scenario leads to a temperature rise of 3 degrees C. Less warming (2 C to 4 C) is projected for the southern Baltic Sea region. There is uncertainty in these projections from different sources (e.g. greenhouse gas emission scenarios, global and regional climate model uncertainties, natural variability). The largest uncertainty is due to the unknown future emissions (RCPs) (Kjellström et al. 2011, Strandberg et al., 2014).

Uncertainties:

Trends in surface temperature can be expected to follow the change in air temperature due to air sea heat exchange. Here global models highly agree at least qualitatively. Under the assumption of moderate greenhouse gas emissions high resolution model simulations project a warming of surface air temperature by up to 8 °C degree in winter and up to 10°C during summer for the Baltic Sea region (Lind and Kjellström, 2008) at the end of the 20th century.

Air temperature extremes:

The strong increase in winter daily mean temperature is most pronounced for the coldest episodes (Kjellström 2004). This is also the case for the most extreme daily maximum and minimum temperatures (Kjellström et al. 2007; Nikulin et al. 2011) with a significant decrease in probabilities of cold temperatures (Benestad 2011). In summer, warm extremes are projected to become more pronounced. For example, Nikulin et al. (2011) showed that warm extremes in today's climate (1961–1990) with a 20-year return value (defined as the temperature that will be exceeded once every 20 years as a statistical average) will occur around once every 5 years in Scandinavia by 2071–2100 according to an ensemble of six RCM simulations, all downscaling GCMs under the SRES A1B scenario.

3.3.1.3 Solar radiation

Mean change is uncertain. Global climate models indicate an increase which is highest over southern Europe and decreases towards north, but still showing a slight increase over the Baltic. However, regional climate model runs could instead show a decrease in surface solar radiation (SSR) over the Baltic area, i.e. there is a large discrepancy in modelled SSR between global and regional models (Bartók et al., 2017). Unknown future aerosol emissions add to the uncertainty.

3.3.1.4 Precipitation

Anna:

Based on fundamentals of the climate system, a warming is expected to result in an intensification of the hydrological cycle i.e. higher precipitation in already wet regions and less precipitation in already dry regions. This is also reproduced by models and scenarios indicate an increase in the entire area during winter and mainly in the northern parts during summer. For some areas a decrease in summer precipitation is expected.

Kommentiert [m20]: Individual scenarios? I suggest to refer to information from IPCC 2013

Kommentiert [JH21]: Missing number

Kommentiert [m22]: Reference?

Erik:

A warmer climate leads to an amplification of the hydrological cycle. For areas with large amounts of precipitation this implies even more precipitation while for dry areas there is a risk of further drying. In the Baltic Sea region this implies increasing
5 precipitation, most notably in the winter half of the year and in the north in summer. There is a large uncertainty as whether precipitation will increase or decrease in summer in the southern part of the Baltic Sea region that is closer to the dry regime of southern Europe.

A warmer future climate will lead to a shortening of the snow season. In winter, climate models project more precipitation and
10 higher temperatures. In many areas this will likely be manifested as more rain and less snow but in some areas that are still cold increasing amounts of snow may be seen.

The degree of change depends on the change in forcing conditions and the regional response of the climate system. Both of these are associated with large uncertainties. In addition, internal natural variability of the climate system adds another level
15 of uncertainty when addressing precipitation changes for a certain time period.

Precipitation extremes:

Anna:

20 More intense precipitation is expected on time scales ranging from single rain showers to long-lasting synoptic-scale precipitation. Where particularly the shorter duration extremes is expected to increase (for example is the return period for 20-year rainfall events on a 1-hour basis decreased to 4 years in model simulations).

Erik:

25 A warmer atmosphere that can hold more water vapor increases the potential for precipitation extremes. Both droughts and heavy rainfall events can become more intense.

Recent climate model simulations performed for other regions in Europe and North America indicate that high-intensity rainfall events associated with summertime convection may generally increase with up to 10-15% per degree of temperature increase.

30 **3.3.1.5 Wind**

Wind:

Wind extremes:

3.3.1.6 Atmospheric chemistry

3.3.2 Land

River runoff, nutrient and carbon loads, river ice, lakes, lake ice, snow, permafrost, terrestrial and freshwater ecosystem,

5 Figures of ensemble mean changes in mean and extreme variables in uncoupled (Ole, Erik) and coupled simulations (Matthias, Christian, Markus)

3.3.2.1 River discharge

Climate change is likely to have a clear influence on the seasonal flow regime as a direct response to changes in the form of the precipitation, as well as by altering the temperature-evapotranspiration regime. (BACC II)

10 For areas in the northern Baltic Sea region presently characterized by spring floods due to snow melt, the floods are likely to occur earlier in the year and their magnitude is likely to decrease owing to less snowfall, shorter snow accumulation period, and repeated melting during winter. As a consequence, sediment transport and the risk of inundation are likely to decrease.

15 In the southern part of the Baltic Sea area, increasing winter precipitation is projected to result in increased river discharge during winter. In addition, groundwater recharge is projected to increase in areas where infiltration capacity is not exceeded currently, resulting in higher groundwater levels. Decreasing precipitation combined with rising temperature and evapotranspiration during summer is projected to result in a drying of the root zone, which would drive increasing irrigation demands in the southern part of the Baltic Sea area. Projections with a process-oriented hydrological model suggest that under RCP 4.5 and RCP 8.5 scenarios the total river flow will increase between 1 and 21% and between 6 and 20%, respectively, 20 illustrating the large uncertainty in hydrological projections (Saraiva et al., 2019). Thereby, the future period 2069-2098 is compared to the reference period 1976-2005 and 4 (3) GCMs are regionalized.

The decrease in snow melt induced spring floods in the southern Baltic rivers contrasts with the situation in the Nordic countries, where changes in winter snowmelt are not yet apparent in the river runoff data, although they are expected in the future (Veijalainen et al. 2010).

25 3.3.2.2 Lakes

3.3.2.3 Terrestrial biogeochemistry

Nutrient loads:

GCMs suggest the north be wetter and the south will be drier (BACC II).

Models suggest land-based nutrient management will have greater effect on loads than uncertainties caused by greenhouse gas emission scenarios (Saraiva et al. 2019)

5 **DOC inputs will increase in areas affected by permafrost thaw (BACC II)**

Existing scenario simulations of the Baltic Sea were carried out with nutrient load scenarios that span the range of plausible future socio-economic conditions from the most optimistic (BSAP) to the worst scenario (Saraiva et al., 2019; Meier et al., 2019).

10 **In the reference scenario, nutrient loads represent the average loads of the period 2010-2012. The high or worst case scenario assumes changes caused by a ‘fossil-fuelled development’ scenario coupled to increasing river runoff. Changes in nitrogen and phosphorus loads were calculated from regional assumptions, e.g., on population growth, changes in agricultural practices such as land and fertilizer use and expansion of sewage water treatment plants (Zandersen et al., 2019).**

3.3.2.4 Terrestrial ecosystems and biodiversity

15

3.3.3 Cryosphere

3.3.3.1 Snow

3.3.3.2 Glaciers

3.3.3.3 Permafrost

20 **3.3.3.4 Sea ice**

In the future it is likely, that the inter-annual variability continues to be very large, although the probability of very strong winters will very likely be lower than in the past. But in climatological sense it is very likely that the maximum sea ice extent of a winter season will decrease. The level ice thickness will also very likely decrease in the future, but there are still larger uncertainties in the thickness of ridged ice. It is likely that the length of the ice season will decrease, but with larger regional differences.

25

There is some indication (low confidence) in climate scenarios that the snow cover on the sea ice will decrease. This will reduce the reflectivity of the surface and more heat can be absorbed and thereby amplifying the ice reduction (Höglund et al., 2017).

3.3.3.5 Cryospheric ecosystems and biodiversity

3.3.4 Ocean and sediment

Water temperature, salinity, stratification, currents, water exchange, sea ice, sea level, wind waves, sediment transports, marine ecosystem (including coastal ecosystems, biogeochemistry), acidification, Figures of ensemble mean changes in mean and extreme variables in uncoupled (Ole, Erik) and coupled simulations (Matthias, Christian, Markus)

3.3.4.1 Water temperature

Water temperature introduction:

The response time of water temperature is short compared to the time scale of climate warming on centennial time scales. In quasi-steady state, the warming at the surface is larger than in the deep water.

With time the heat spreads downward through different processes and eventually the whole water column warms up. This has consequences for the stratification, eutrophication and sea level rise.

Water temperature: The temperature of the oceans are rising [high confidence] (e.g. Balmaseda et al., 2013a, 2013b) and will continue to rise [medium confidence]. Regional scenarios for the Baltic Sea project an increase of SST of 1 C (RCP2.6) to 4 C (RCP8.5) in 2100 relative to the period 1970 to 1999. The projections include a possible range of 1.5 C (RCP2.6) and 2.5 C (RCP8.5). The SST changes in the RCP8.5 scenarios are significantly above the natural variability (Source SMHI).

Water temperature extremes: In the RCP4.5 and RCP8.5 scenarios an increase in the occurrence of tropical nights over the Baltic Sea has been observed. Tropical nights are an indicator for heat waves. This has consequences for the water temperature which will reach record breaking values from year to year much more often under projected climate change in RCP8.5 (Meier et al., 2019).

RCP 8.5: the sea surface temperature response in the Baltic Sea varies between + 2.5 and + 4.7 K depending on the applied global model scenario (Gröger et al., 2019)

3.3.4.2 Salinity and saltwater inflows

Future changes in salinity are difficult to project as they depend on the large-scale atmospheric moisture transport to the Baltic Sea drainage basin (Schinke and Matthäus, 1998), local wind fields (Lass and Matthäus, 1996) and global mean sea level

(GMSL) rise (Hordoir et al., 2017; Meier et al., 2017). However, to robustly estimate future moisture transport and runoff to the Baltic Sea and regional wind fields requires high resolution atmosphere models to resolve local orography along continental water sheds. Furthermore, even on a global scale projections of the water cycle and sea level rise suffer from uncertainties (IPCC, 2013). This makes projections for the future development of salinity highly uncertain (BACC II Author team, 2015).

5

Due to the projected increased freshwater supply from the catchment area by about 1 to 21% at the end of the century depending on the climate model, surface and bottom salinity is projected to decrease by about 0.6 g kg^{-1} in the ensemble mean with a large spread among the ensemble members (Saraiva et al., 2019). However, ensemble studies taking rising global mean sea level and changes in wind fields and river flows into account do not exist (Meier et al., 2018b).

10

Hordoir et al. (2017) investigated the influence of rising GMSL on saltwater inflows into the Baltic Sea. They performed idealized model sensitivity experiments using a regional ocean general circulation model covering the North Sea and the Baltic Sea. Hordoir et al. (2017) found a non-linear increase in saltwater inflow intensity and frequency with rising GMSL. However, their explanation of reduced mixing in the Danish straits was shown to be wrong (Arneborg, 2016). Instead, Arneborg (2016) proposed an alternative theory. Due to the smaller depth, the volume flux through the Sound is more sensitive to GMSL rise than that through the Belt Sea. Under present conditions, the amount of dense water passing the Drogden sill in the Sound is determined by a baroclinic control in the narrow northern end of the Sound (Nielsen, 2001). With rising GMSL this control is degraded and relatively more saltwater is transported into the Baltic Sea compared to the expected increase when the transport change is proportional to the area of the limiting cross section.

15

Assuming a negligible impact of GMSL rise, the intensity and frequency of MBIs were projected to remain unchanged, with a potential tendency of a slight increase (Schimanke et al., 2014). However, in future high-end global mean sea level projections, reinforced saltwater inflows result in higher salinity and increased vertical stratification compared to present conditions (Meier et al., 2017; Saraiva et al., 2019b).

25 3.3.4.3 Stratification

Theoretical considerations imply that increased freshwater supply (rain, snow melt) over the Baltic Sea drainage basin accompanied by the supply of deep salt rich waters from the North Sea as well as warming of the surface layer would favor stronger stratification. Thus the future development of stratification mainly depends on how the Baltic Sea surface will warm up compared to deeper layers and how the freshwater supply will change (which in turn is linked to atmospheric moisture transport to the region).

30

The complex interplay between temperature change, wind change and changing precipitation makes it difficult to predict future climate effects on stratification. Furthermore, stratification is subject to complex processes like winter convection which are likely to change with future climate change.

Most model studies project increasing sea surface temperatures for the end of the 21st century for the Baltic Sea (BACC II). Future climate projections for the Baltic Sea based on previous assumptions on greenhouse emission scenarios imply an SST increase between 2-4 °C (Meier 2015). More recent scenarios with revised emission pathway imply an increase of 2-3 °C (Saraiva et al., 2018).

Most studies suggest rather moderate changes or slightly lowered surface salinities in the Baltic Sea (Meier, 2015). In certain weaker stratified regions like the Gulf of Finland or Bothnian Bay changes in deep salinity could lead to decreased stratification (Meier 2015)

Little is known about future trends in salt inflows to the Baltic Sea which mainly impact on deep water salinity (with positive effect on stratification). Favorable atmospheric preconditions i.e. the prevalent wind regime for salt inflows have been reported to occur slightly more frequent in future (Schimanke et al., 2014).

Few available studies suggest a tendency towards only slightly modified mean and extreme wind speed over the Baltic Sea at the end of the century (e.g. Kjellström et al., 2011).

The above described changes in governing parameters would favor stratification to increase. In fact few studies report that stratification is likely to increase in future climate (Gröger et al., 2019).

The entire ensemble consistently indicates a basin-wide intensification of the pycnocline (9–35%) for the Baltic Sea and a shallowing of the pycnocline depth in most regions as well (Gröger et al., 2019).

3.3.4.4 Overturning circulation

All ensemble members indicate a strengthening of the zonal, wind driven near surface overturning circulation in the southwestern Baltic Sea towards the end of the twenty-first century whereas the more thermohaline driven overturning at depth is reduced by ~ 25% (Gröger et al. 2019). In the Baltic proper, the meridional overturning shows no clear climate change signal. However, three out of five ensemble members indicate at least a northward expansion of the main overturning cell. In the Bothnian Sea, all ensemble members show a significant weakening of the meridional overturning.

3.3.4.5 Sea level

Since the ice has melted that covered Scandinavia during the last ice age the earth crust is rebounding. The highest rates of land uplift of 10 mm/year are found at Höga Kusten in the Bothnian Bay. The land uplift diminishes away from this center and is very small along the German and Polish coasts of the Baltic Sea. This process is expected to continue for thousands of years [high confidence]. Along most of the Baltic Sea coasts the land uplift causes sea level to fall. On top of this change, global mean sea level rise will continue to raise sea levels in the Baltic Sea at an increasing rate [high confidence] (Church et al., 2013). When the rate of sea level rise becomes larger than the land uplift, sea level will start to rise relative to land. First in the southern Baltic Sea and with accelerating rates of global mean sea level rise the line of rising sea level will move northward.

Kommentiert [m23]: Meier and Saraiva (2019). Would be good to specify the underlying scenarios? A1B, A2 or RCP 4.5 and 8.5. 2-3°C corresponds to ensemble mean changes under RCP 4.5 and 8.5 averaged for the entire Baltic Sea.

Kommentiert [JH24]: What time of year? Annual mean? Link to sea ice cover? Which region does this refer to (in the Baltic)?

Kommentiert [JH25]: Why do we present an older projection if we have a newer more specific one?

Kommentiert [m26]: Merge with the other variables

Kommentiert [m27]: But not with the same rising rate? Replace with hundred years

During this century melting ice sheets in Antarctica and Greenland will contribute more to the total than in the past [medium confidence] (Church et al., 2013). The sea level rise due to melting ice is not distributed uniformly around the global ocean. Sea level rise from melting ice sheets in Antarctica is more pronounced in the northern hemisphere because the missing ice mass has a smaller gravitational pull on the surrounding water. As in Scandinavia, the missing ice causes the earth crust to rebound which makes the sea level rise slower where the ice is melting [medium confidence]. Based on these processes the mean sea level rise in the Baltic Sea is projected to amount to 80% of the global mean sea level rise (Grinsted, 2015). Recent efforts since the latest IPCC report that focused on the contribution of Antarctic ice sheets to global mean sea level rise have shown that the interaction of warming ocean water, melting the ice sheets from below can lead to instabilities in the ice sheet dynamics. The ice sheets flowing from land into the ocean are in contact with the ocean floor out to the grounding line. From there on outward the ocean is melting the ice from below and the ice sheets become thinner and lighter. If the weight of the ice sheet becomes less than the weight of the ocean water it replaces, it floats up and away. The grounding line retreats inland where the ice sheet is thicker and the ice flow larger and reinforces the ice loss (Mercer, 1978). This and related feedback loops could lead to an extra meter of sea level rise until the end of the century [low confidence] (e.g. Sweet et al., 2017). The most recent estimates (Bamber et al., June 2019) for global mean sea level rise in 2100 relative to 2000, including these potential contributions (including land water storage) are 69 cm and 111 cm for low and high sea level scenarios, respectively. For the high sea level scenario the likely range (5% to 95%) is between 62 cm and 238 cm.

Kommentiert [JH28]: Are there studies for the Baltic Sea that tried to combine all the effects? Meier et al. 2004

Kommentiert [m29]: Baltic Sea numbers?

Sea level extremes:

In the Baltic Sea, return levels show a low to moderate increase for increasing return periods (HazardSupport, 2018)(true everywhere?). The 100-year return level for Stockholm for example is estimated at 102 cm, while the 200-year return level is 107 cm. Both estimates have a large uncertainty of 25 cm and 30 cm, respectively (SMHI report regeringsuppdraget). Even under moderate increase of mean sea level in the Baltic Sea, extreme events that are rare today will be much more common towards the end of the century. In the above example a sea level rise of 20 cm in 2100 will turn the 100-year storm surge in Stockholm into an event that occurs every 10 years on average [medium confidence]. Thus, the main driver of changes in Baltic Sea storm surges is the global mean sea level rise [medium confidence].

3.3.4.6 Waves

3.3.4.7 Sedimentation

If the relative sea level rise in the southern Baltic Sea follows the mean value of RCP2.6 projection, which is ~0.24 m until 2065, the rate of sea level rise in this region (~1.2 mm/yr for the past few decades) would be accelerated. As a natural consequence coastal erosion would be regionally enhanced to fill the increased underwater accommodation space. The extent of enhancement in erosion depends on not only the sea level but also storms.

Due to the impact of the prevailing westerly winds, the dominant sediment transport will remain eastwards along a major part of the southern Baltic coast, but with a high variability along some local coast sections which have small incidence angle of incoming wind-waves.

5

Development of the foredunes will continue in prograding coasts.

A critical threshold, which distinguishes a linear and a non-linear (following a quadratic or a higher power law) relationship between foredune height and rate of relative sea level rise, seems to exist in the southern Baltic Sea coast. If a rise by 0.3 m in the relative sea level (RCP8.5) would occur by 2065 in this region, such critical threshold would probably be reached before 2050, causing drastic change on the foredune characteristics and much stronger erosion on cliffs and old coastal dunes.

10

3.3.4.8 Marine carbonate and biogeochemistry chemistry

Carbonate chemistry (incl. air-sea exchange of CO₂) and acidification

15

The atmospheric CO₂ concentration will increase in the future and influence the marine CO₂ system.

Kuznetsov et al. (2013)

Omstedt et al. (2012)

20

Marine biogeochemistry

Oxygen concentration and hypoxia

Projected warming and global mean sea level rise may reinforce eutrophication and oxygen depletion in the Baltic Sea by reducing air-sea and vertical transports of oxygen -, intensifying internal nutrient cycling, and increasing river-borne nutrient loads. However, the response of deep-water oxygen conditions to changing climate will mainly depend on the nutrient load scenario. In the case of high (low) nutrient loads, the impact of the changing climate would be considerable (negligible). Scenario simulations suggest that the complete implementation of the Baltic Sea Action Plan (BSAP) resulting in required load reductions will lead to a significantly improved ecosystem state of the Baltic Sea irrespective of the driving global climate model. The latter was shown for the ecosystem indicators water clarity and summer mean oxygen deficit due to biogeochemical oxygen consumption compared to saturated oxygen conditions.

30

For the end of the century (2069-2098), hypoxic area is projected to change only slightly in the ensemble mean under reference/present-day (-14 ... -5%) and high (-2 ... +5%) nutrient load scenarios compared to the period 1976-2005.

5 Under the BSAP scenario, hypoxic area will be reduced by 50 to 60% in the ensemble mean at the end of the century compared to 1976-2005.

3.3.4.9 Marine ecosystems and biodiversity

Pelagic habitats (incl. phytoplankton and zooplankton community structure, spring blooms, functional traits etc.)

10 Spring blooms start earlier and are longer

Both phyto- and zooplankton communities are composed of smaller species

Increasing cyanobacteria blooms (why?)

Shifting diatom/dinoflagellate ratios

15 **Benthic habitats (incl. benthic organisms and community structure, functional traits)**

Fish

Bauer et al. (2018, 2019)

20 **Water birds**

Marine mammals

Non indigenous species

Ecotoxicology

Ecosystem function

25

4 Interactions with other drivers

Summary of the BEAR paper by Reckermann et al.:

Marine (shipping), terrestrial and airborne traffic

30 Tourism

Built structures (incl. offshore renewable energy constructions and maintenance)

Fisheries

Aquaculture

Eutrophication

Pollution and hazardous substances

5 Coastal defence and erosion

Marine Protected Areas (MPAs)

Forestry

Urbanisation and population changes

Aerosols

10 Land cover changes

Etc.

5 Comparison with other coastal sea regions

6 Knowledge gaps

15 **Large-scale circulation:**

NAO:

While CMIP5 climate models are able to simulate the main features of the NAO, its future changes might be sensitive to boundary processes, which are not yet well represented in many climate models (IPCC, 2013).

Blocking:

20 Most CMIP5 models still underestimate the frequency of blocking over the Euro-Atlantic sector (IPCC, 2013).

AMO:

Since the observational record is relatively short, our understanding of the AMO and its possible changes largely depend on models, whose assumptions are difficult to verify (Knight, 2009).

25 However, while possible changes in these climate phenomena contribute to the uncertainty in near-term climate projections, they are not the main driver of the projected warming over Europe by the end of the century (Cattiaux et al., 2013; IPCC, 2013).

Air temperature:

30 There are limitations in the knowledge concerning the link to changes in large-scale circulation patterns.

Kommentiert [m30]: More specific

Kommentiert [m31]: More specific

Surface solar radiation:

Multidecadal variations in SSR are generally not well captured by current climate model simulations (Allen et al. 2013, Storelvmo et al. 2018). The extent to which the observed SSR variations are caused by natural variation in cloudiness induced by atmospheric dynamic variability (Stanhill et al. 2014, Parding et al. 2016), or anthropogenic aerosol emissions (Ruckstuhl et al. 2008, Philipona et al 2009, Wild 2012, Storelvmo et al. 2018), or perhaps additional causes, is not well quantified.

Precipitation:

Focus on different aspects of precipitation characteristics, different methods and different data sets used in various national studies in the Baltic Sea region implies that the picture of the precipitation climate including its past changes is not fully coherent.

Even if climate scenarios are becoming more frequent and there is now a growing ensemble of relatively high-resolution regional climate scenarios for Europe the scenarios still only samples a subset of the global climate model projections assessed by the IPCC. This means that the uncertainties of future climate change in the Baltic Sea region is not fully captured at adequate horizontal resolution for pursuing detailed studies of climate change impacts in the region.

New, very high-resolution so called convective-permitting climate models operating at grid spacing of 1-3 km are lacking for the Baltic Sea region. In other regions such models have shown better agreement with observations in representing precipitation extremes and sometimes also a larger climate change signal compared to the more traditional "high-resolution" models operating at c. 10 km grid spacing.

In addition: Analysis of changes in extreme precipitation is uncertain due to the problem with statistical significance of changes in extreme events.

Runoff:

The impact of how climate model results are transferred to the hydrological model are still inadequately understood. More research is needed to quantify the accuracy and uncertainty associated with various bias correction methods. Several uncertainties are associated with impact modelling, including parameter uncertainty and model structure uncertainty. The values of the parameters of a hydrological model are normally found through calibration against historical data and are always associated with uncertainty. This uncertainty will translate into uncertainty in the projected changes.

Nutrient loads:

How fertilization practices, crops grown, and land use will change in response to climate change.

Salinity:

Due to uncertain changes in regional water cycles (precipitation) and global sea levels, the confidence in future salinity projections is low (BACC II Author Team, 2015; Meier et al., 2019). Ensemble studies taking rising global mean sea level and changes in wind fields and river flows into account do not exist but would be needed (Meier et al., 2018).

Sea level:

There is potential for mean sea level rise in the Baltic Sea that is caused by the freshening of the Baltic Sea. Under a warming climate precipitation patterns over the Baltic Sea drainage basin change (EN-CLIME precipitation) and the river discharge is expected to increase (EN-CLIME discharge). Fresher water in the Baltic Sea will take up more volume and sea levels in the Baltic Sea will rise. There has been no comprehensive study so far to assess the impact of halosteric contributions to the Baltic Sea level. Also the impact of changing land water storage for the Baltic Sea region on Baltic Sea levels needs to be assessed (Haasnoot et al., 2015 for the Netherlands).

On the global scale it has been shown (Bingham and Hughes, 2012) that sea level will raise proportionally more on the shallow shelf regions around the continents than in the deeper, open ocean. For the Baltic Sea no such study exists that could answer whether mean sea level rise applies to coastal regions in the same way as for the open Baltic Sea.

Storm surges and other hazards can turn into disaster if they occur concurrently. There have been some examples and studies (e.g. Zscheischler et al., 2018) that have shown that the impact of multiple hazards can be much worse than a single extreme event. Not much is known today about the interaction of extreme events with a focus on storm surges.

The coverage in time and space of sea level measurements in the Baltic Sea is among the best in the world. Many records date back to late 19th century. Nevertheless, the observational record tends to underestimate the natural variability that is inherent in storm floods (Lang and Mikolajewicz, 2019) in the Baltic Sea. One example is the Backafloeden 1872. It is not clear which return period should be attributed to this event (Fredriksson, 2016, 2017). With today's observational time series it cannot be categorized. Two conclusions can be drawn here. First, more research is needed to investigate the influence of natural variability on storm surges in the Baltic Sea. Second, long measurement records are of vital importance to assess natural variability in coastal flooding around the Baltic Sea.

Sea level rise for coastal cities around the Baltic Sea is more sensitive to the thermosteric expansion of the global ocean compared to coastal cities around the coasts of the Atlantic or Pacific Ocean (Larour et al., 2017). Also, ice sheet melt in Antarctica has a disproportional high impact in the Baltic Sea while ice sheet melt in Greenland has no effect. Sea level research focused on the Baltic Sea may take this into account.

Kommentiert [m32]: Why? And why is there a policy gap related to this fact?

Kommentiert [m33]: Missing reference

Sedimentation:

We lack a comprehensive understanding of alongshore sediment transport and associated spatial and temporal variability along the Baltic coast. In general, an eastward transport dominates along a major part of the southern Baltic coast due to the impact of the prevailing westerly winds. However, the intensity of secondary transport induced by easterly and northerly winds is much less understood. Its combination with storm surges further complicated the understanding because in such circumstance the sandy dunes and cliffs are exposed to highest erosional impact.

Another knowledge gap in understanding coastal erosion in response to future climate change is on the impact of water levels and the submergence of the beach.

Anthropogenic influence imposes one of the largest uncertainty in sediment transport and coastal erosion.

An engineering structure (e.g. pier, seawall) influences coastline change at a much larger spatial scale than the dimension of the structure itself.

Changes in carbonate chemistry (incl. air-sea exchange of CO₂) and acidification:

Due to the high spatial and temporal variability of the seawater pCO₂ it is not known whether the Baltic Sea as a whole is a net sink or net source of CO₂.

It is unclear what is the source of total alkalinity increase in the Baltic Sea, and whether it will continue in the future with the same magnitude.

Oxygen concentrations:

A recent assessment suggests that the biggest uncertainties in projections of biogeochemical cycles are caused by (1) poorly known current and future bioavailable nutrient loads from land and atmosphere, (2) the setup of numerical scenario experiments (including the spin up strategy), (3) differences between the projections of global and regional climate models, in particular, with respect to the global mean sea level rise and regional water cycle, (4) differing model-specific responses of the simulated biogeochemical cycles to long-term changes in external nutrient loads and climate of the Baltic Sea region, and (5) unknown future greenhouse gas emissions.

Pelagic habitats:

Long-term development, effect on species composition and carbon transfer (including settling)

Evaluation on the importance of the climate change component
 Evaluation on the importance of the climate change component
 The connection still not clear, the effect may depend on other factors

5

6.1 Knowledge gaps - consensus

6.2 Knowledge gaps - dissensus

6.3 Research needs [this is what we need to know]

10 **7 Key messages**

Confirmed and new knowledge of BACC III

Author contributions

| Chapter | Title | Authors | Pages ¹ |
|---------|--|--|--------------------|
| 1 | Introduction | All | 8 |
| 1.1 | Overview | H.E.M. Meier and all | 0.5 |
| 1.2 | Baltic Sea Region characteristics | H.E.M. Meier, E. Kjellström, J. Käyhko and all | 2 |
| 1.3 | Global climate change | E. Kjellström and M. Gröger | 2 |
| 1.4 | The BACC process | M. Reckermann | 0.5 |
| 1.5 | Summary of BACC I and BACC II key messages | H.E.M. Meier and all | 3 |
| 1.5.1 | Past climate change | H.E.M. Meier and all | 1 |
| 1.5.2 | Recent climate change | H.E.M. Meier and all | 1 |
| 1.5.3 | Future climate change | H.E.M. Meier and all | 1 |

¹ Without references, figures and tables (just approximations)

| | | | |
|---------|---|--|-----|
| 2 | Methods | H.E.M. Meier, C. Dieterich, O. Bøssing-Christensen, E. Kjellström | 1.5 |
| 2.1 | Assessment of literature | H.E.M. Meier and M. Reckermann | 0.5 |
| 2.2 | Climate model data | C. Dieterich, E. Kjellström and O. Bøssing-Christensen | 0.5 |
| 2.3 | Uncertainty estimates | E. Kjellström and H.E.M. Meier | 0.5 |
| 3 | Current state of knowledge | All | 45 |
| 3.1 | Past climate change | E. Zorita | 5 |
| 3.2 | Present climate change | All | 20 |
| 3.2.1 | Present climate change - Atmosphere | A. Rutgersson, M. Stendel, C. Frauen, M. Quante | 6 |
| 3.2.1.1 | Large-scale circulation | C. Frauen | |
| 3.2.1.2 | Air temperature | A. Rutgersson | |
| 3.2.1.3 | Solar radiation | A. Rutgersson | |
| 3.2.1.4 | Precipitation | M. Stendel | |
| 3.2.1.5 | Wind | M. Stendel | |
| 3.2.1.6 | Atmospheric chemistry | M. Quante | |
| 3.2.2 | Present climate change - Land | J. Käyhko | 4 |
| 3.2.2.1 | River discharge | J. Käyhko | |
| 3.2.2.2 | Lakes | J. Käyhko | |
| 3.2.2.3 | Terrestrial biogeochemistry | C. Humborg | |
| 3.2.2.4 | Terrestrial ecosystems and biodiversity | P. Miller | |
| 3.2.3 | Present climate change - Cryosphere | J.J. Haapala | 4 |
| 3.2.3.1 | Snow | S. Rasmus | |
| 3.2.3.2 | Glaciers | S. Rasmus | |
| 3.2.3.3 | Permafrost | S. Rasmus | |
| 3.2.3.4 | Sea ice | J.J. Haapala | |
| 3.2.3.5 | Cryospheric ecosystems and biodiversity | | |
| 3.2.4 | Present climate change – Ocean and sediment | H.E.M. Meier, M. Kniebusch, B. Hünicke, E. Zorita, M. Viitasalu, K. Kulinski, W. Zhang | 6 |

| | | | |
|---------|---|---|----|
| 3.2.4.1 | Water temperature | M. Kniebusch, H.E.M. Meier | |
| 3.2.4.2 | Salinity and saltwater inflows | M. Kniebusch, H.E.M. Meier | |
| 3.2.4.3 | Stratification | M. Kniebusch, H.E.M. Meier | |
| 3.2.4.4 | Overturning circulation | H.E.M. Meier | |
| 3.2.4.5 | Sea level | B. Hünicke, E. Zorita, C. Dieterich | |
| 3.2.4.6 | Waves | B. Hünicke | |
| 3.2.4.7 | Sedimentation | W. Zhang | |
| 3.2.6.8 | Marine carbonate and biogeochemistry | K. Kulinski and H.E.M. Meier | |
| 3.2.6.9 | Marine ecosystems and biodiversity | M. Viitasalu | |
| 3.3 | Future climate change | All | 20 |
| 3.3.1 | Future climate change - Atmosphere | O. Bøssing-Christensen, E. Kjellström, M. Quante | 6 |
| 3.3.1.1 | Large-scale circulation | O. Bøssing-Christensen, E. Kjellström | |
| 3.3.1.2 | Air temperature | O. Bøssing-Christensen, E. Kjellström | |
| 3.3.1.3 | Solar radiation | O. Bøssing-Christensen, E. Kjellström | |
| 3.3.1.4 | Precipitation | O. Bøssing-Christensen, E. Kjellström | |
| 3.3.1.5 | Wind | O. Bøssing-Christensen, E. Kjellström | |
| 3.3.1.6 | Atmospheric chemistry | M. Quante | |
| 3.3.2 | Future climate change - Land | ???, J. Käyhko, O. Bøssing-Christensen, E. Kjellström, C. Dieterich | 4 |
| 3.3.2.1 | River discharge | J. Käyhko | |
| 3.3.2.2 | Lakes | J. Käyhko | |
| 3.3.2.3 | Terrestrial biogeochemistry | C. Humborg | |
| 3.3.2.4 | Terrestrial ecosystems and biodiversity | P. Miller | |
| 3.3.3 | Future climate change - Cryosphere | J.J. Haapala | 4 |
| 3.3.3.1 | Snow | O. Bøssing-Christensen, E. Kjellström | |
| 3.3.3.2 | Glaciers | O. Bøssing-Christensen, E. Kjellström | |
| 3.3.3.3 | Permafrost | O. Bøssing-Christensen, E. Kjellström | |
| 3.3.3.4 | Sea ice | J.J. Haapala | |

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|--------------------|--|---|---|
| 3.3.3.5 | Cryospheric ecosystems and biodiversity | | |
| 3.3.4 | Future climate change – Ocean and sediment | H.E.M. Meier, B. Hünicke, E. Zorita, M. Viitasalu, K. Kulinski, W. Zhang, C. Dieterich, M. Gröger | 6 |
| 3.3.4.1 | Water temperature | H.E.M. Meier, C. Dieterich, M. Gröger | |
| 3.3.4.2 | Salinity and saltwater inflows | H.E.M. Meier, C. Dieterich, M. Gröger | |
| 3.3.4.3 | Stratification | M. Gröger, C. Dieterich, H.E.M. Meier | |
| 3.3.4.4 | Overtuning circulation | M. Gröger, C. Dieterich, H.E.M. Meier | |
| 3.3.4.5 | Sea level | C. Dieterich, B. Hünicke, E. Zorita | |
| 3.3.4.6 | Waves | C. Dieterich, B. Hünicke, | |
| 3.3.4.7 | Sedimentation | W. Zhang | |
| 3.3.4.8 | Marine carbonate and biogeochemistry | K. Kulinski, H.E.M. Meier | |
| 3.3.4.9 | Marine ecosystems and biodiversity | M. Viitasalu | |
| 4 | Interaction with other drivers | M. Reckermann | 1 |
| 5 | Comparison with other coastal sea regions | M. Quante, M. Reckermann, H.E.M. Meier, E. Kjellström, J. Käyhko | 2 |
| 6 | Knowledge gaps | All | 2 |
| 7 | Key messages | All | 2 |
| Figures and Tables | Analysis of observed time series | M. Kniebusch | |
| Figures and Tables | Analysis of scenario simulations | O. Bøssing-Christensen, E. Kjellström, H.E.M. Meier, F. Börgel, C. Dieterich, M. Gröger, G. Väli | |

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Kommentiert [m34]: In red relevant references after BACC II

Kommentiert [m35]: Reference format of ESD

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Tables (Ole, Erik, Madline, Florian, Germo, Christian, Matthias, Markus and all)

- 5 **Table 1.** Ensemble mean changes in sea surface temperature (SST) (in °C) in ECOSUPPORT, BalticAPP RCP 4.5 and BalticAPP RCP 8.5 scenario simulations averaged for the Baltic Sea including the Kattegat between 1978-2007 and 2069-2098. (DJF = December, January, February, MAM = March, April, May, JJA = June, July, August, SON = September, October, November)

| Δ SST | DJF | MAM | JJA | SON | Annual mean |
|------------------------|-----|-----|-----|-----|-------------|
| ECOSUPPORT SRES A1B | 2.5 | 2.8 | 2.8 | 2.5 | 2.6 |
| BalticAPP RCP 4.5 | 1.7 | 1.9 | 2.0 | 1.8 | 1.8 |
| BalticAPP RCP 8.5 | 2.9 | 3.2 | 3.3 | 3.0 | 3.1 |

- 10 **Table 2.** Ensemble mean changes in annual mean sea surface salinity (SSS) (in g kg⁻¹), annual mean bottom salinity (BS) (in g kg⁻¹) and winter mean sea level (SL) (in cm) in ECOSUPPORT, BalticAPP RCP 4.5 and BalticAPP RCP 8.5 scenario simulations averaged for the Baltic Sea including the Kattegat between 1978-2007 and 2069-2098.

| Annual changes | ECOSUPPORT | BalticAPP RCP 4.5 | BalticAPP RCP 8.5 |
|----------------|------------|----------------------|----------------------|
| Δ SSS | -1.5 | -0.7 | -0.6 |
| Δ BS | -1.6 | -0.6 | -0.6 |
| Δ SL | 5.5 | 0.4 | 3.7 |

Table 1 Climate change (2070–2099 minus 1970–1999) response of the Baltic Sea calculated for different environmental parameters

| | MPI-ESM | HadGCM | EC-Earth | GFDL | IPSL |
|-------------------------------------|---------|--------|----------|-------|-------|
| Δ SST (K) | +2.27 | +4.67 | +3.70 | +2.50 | +3.52 |
| Δ SSS (g kg^{-1}) | -1.47 | -1.05 | -0.59 | -1.90 | -2.27 |
| Pycnocline intensity (%) | +9.5 | +35 | +25 | +15 | +21 |
| Stratified area (%) | +23 | +78 | +100 | +36 | +53 |

HadGCM delta SSS = -1 g/kg

Ensemble mean SSS changes (Gröger et al., 2019): -1.8 g kg⁻¹

Figures (Ole, Erik, Madline, Florian, Germa, Christian, Matthias, Markus and all)

5 Preliminary list of figures

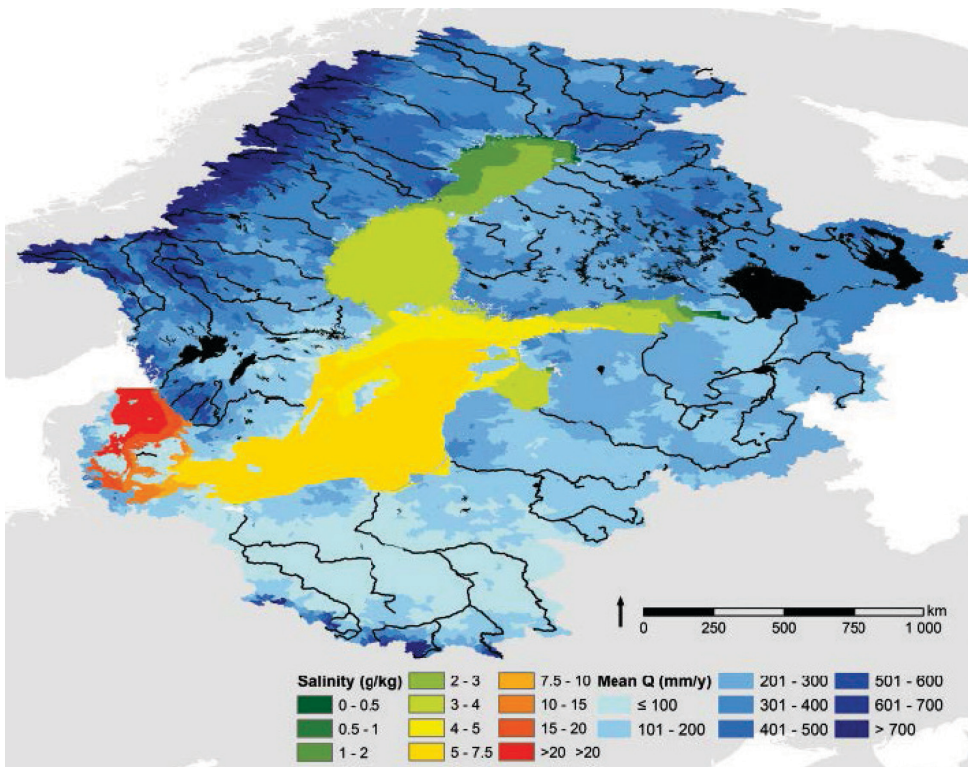


Figure 1: Baltic Sea and catchment area. (Source: Meier et al., 2014)

5 Past climate

....

Present climate

1. NAO
2. AMO
3. Blocking index/Scandinavia
4. Air temperature Stockholm 1756-2018
5. Annual and seasonal mean surface air temperature over the Baltic Sea calculated from CRU 1871-2018

6. Anomaly time series of annual and seasonal precipitation over Sweden, 1860–2004 (reference period 1961–1990) (somebody at SMHI who is updating the time series?) (BACC I Author Team, 2008)
7. Precipitation CRU ???
8. Runoff at single stations with homogeneous data ???
- 5 9. Runoff data for the entire Baltic Sea from SMHI (Göran Lindström), see BACC II Author Team (2015), Fig. 5.18 (I can ask him if he has time to update the figure)
10. Sea surface and bottom water temperature at selected stations (with sufficient data), summer and annual mean with data corrected for seasonal biases (following Hagen Radtke), since 1900
11. Sea surface and bottom salinity at selected stations (with sufficient data), summer and annual mean with data corrected for seasonal biases (following Hagen Radtke), since 1900
- 10 12. Temperature at MARNET station data Darss Sill, Arkona Basin, Hagen curve (Volker Mohrholz)
13. Salinity at MARNET station data Darss Sill, Arkona Basin
14. Maximum annual ice extent 1720/1721-2018/2019, 1900-2018 (see Table 8.1, Fig. 8.3, etc., BACC II)
15. Ice thickness at Kemi, Loviisa
- 15 16. Duration of the ice season
17. Ice breakup Tallinn ???
18. Swedish sea level climate indicator (Hammarklint T, 2009: Swedish Sea Level Series – A Climate Indicator. Swedish Meteorological and Hydrological Institute), already available for 2018
19. Sea level Stockholm 1774-2018

20

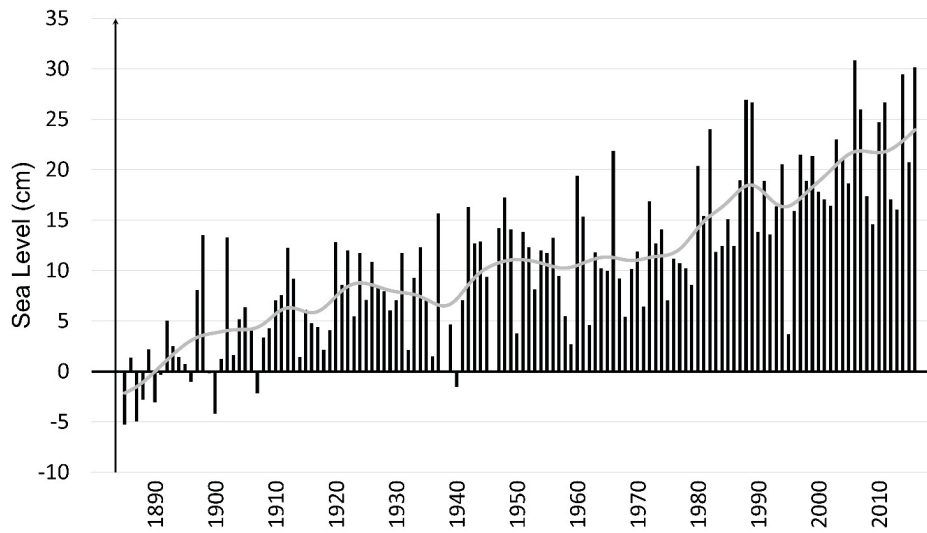


Figure 2: Annual mean sea level changes in centimeters for 14 Swedish mareographs since 1886. The data are corrected for land uplift. The grey line shows a smoothed curve. (Source: SMHI)

5

Future climate

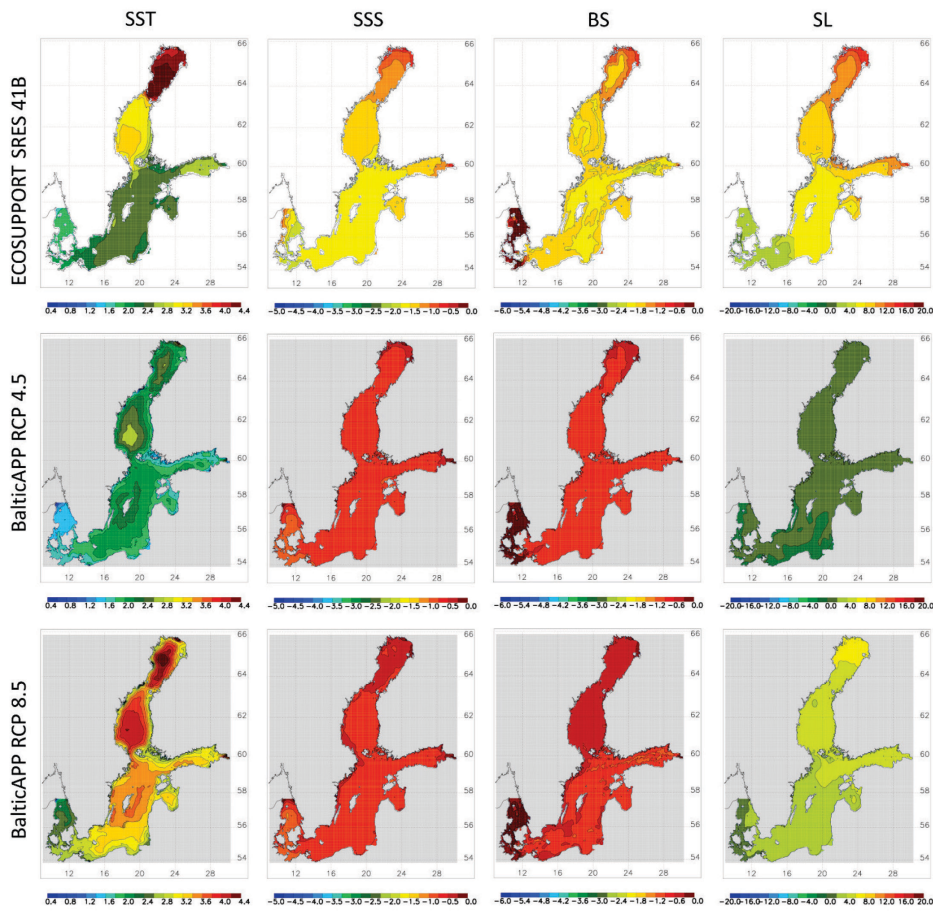


Figure 3: From left to right changes of summer (June – August) mean sea surface temperature (SST) ($^{\circ}\text{C}$), annual mean sea surface salinity (SSS) (g kg^{-1}), annual mean bottom salinity (BS) (g kg^{-1}), and winter (December – February) mean sea level (SL) (cm) between 1978-2007 and 2069-2098 are shown. From top to bottom results of the ensembles ECOSUPPORT (white background), BalticAPP RCP 4.5 (grey background) and BalticAPP RCP 8.5 (grey background) are depicted.

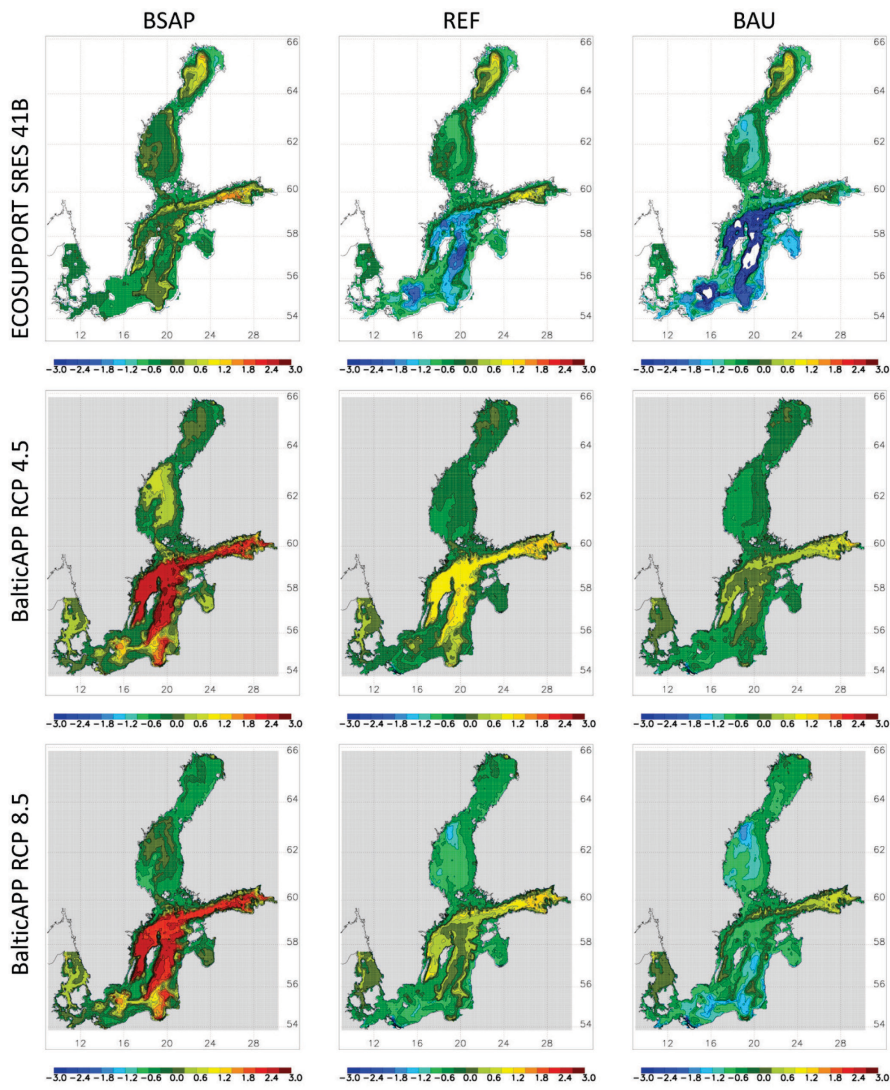


Figure 4: Ensemble mean summer (June – August) bottom dissolved oxygen concentration changes (mL L^{-1}) between 1978-2007 and 2069-2098. From left to right results of the nutrient load scenarios Baltic Sea Action Plan (BSAP), Reference (REF) and Business-As-Usual (BAU) are shown. From top to bottom results of the ensembles ECOSUPPORT (white background), BalticAPP RCP 4.5 (grey background) and BalticAPP RCP 8.5 (grey background) are depicted.

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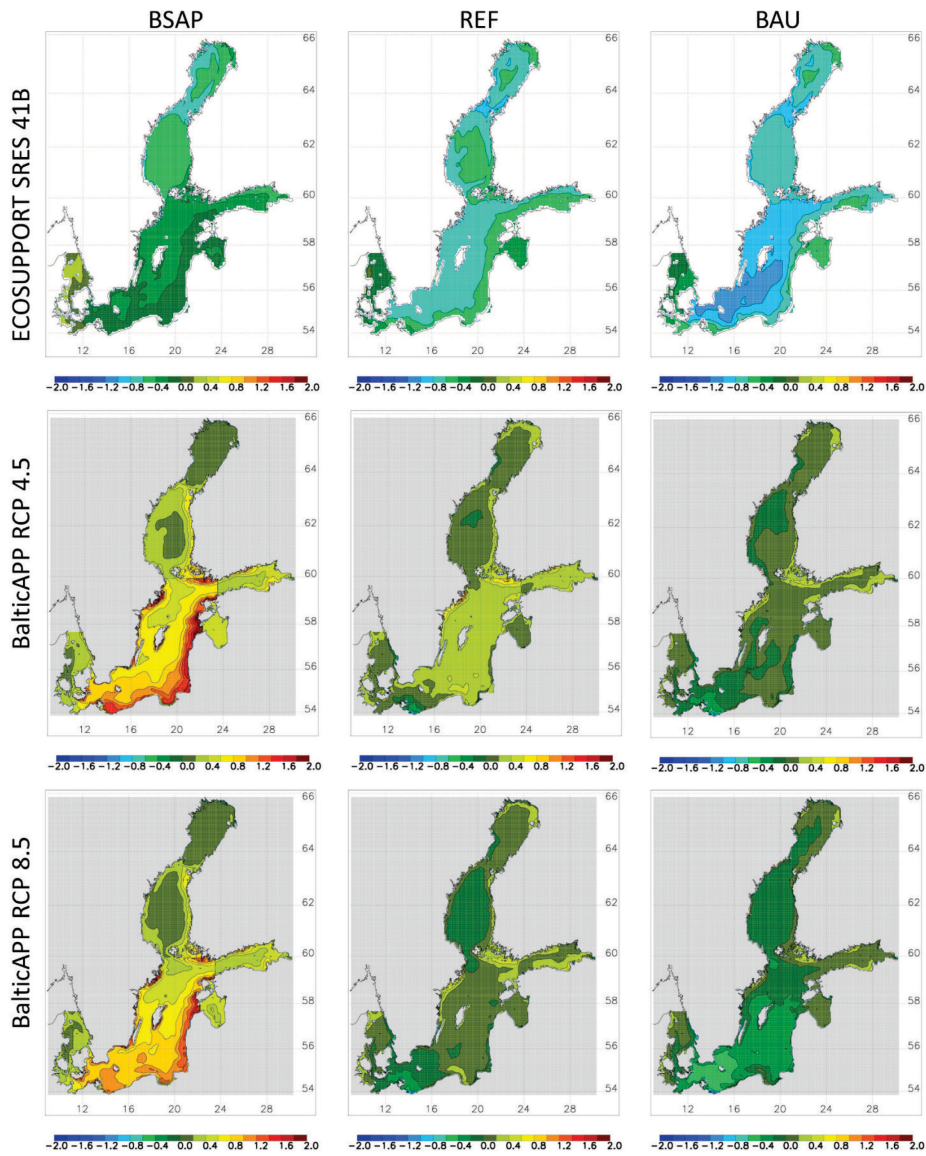


Figure 5: As Fig. 4 but for annual mean Secchi depth changes (m).

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